Hydrogeophysics Exploration Approximately Ain Al Atti (Region of Southeast Erfoud Morocco)

¹Dakir Ibrahim, ²Benamara Ahmed, ¹Aassoumi Habiba, ³Ouallali Abdessalam and ⁴Ait Bahammou Youssef

 ¹Laboratory of Environmental Geology and Natural Resources, Cartography and Digital Technologies Team, Department of Geology, Faculty of Science, Abdelmalek Essaadi University, Tetouan. Morocco
²Laboratory of Applied Geophysics and Modeling in Geoenvironment, Department of Geology, National School of Arts and Crafts, Moulay Ismail University 50500 Meknes, Morocco
³Laboratory of Engineering and Environment, Faculty of Sciences and Technics, Hassan II University.28810, Mohammadia, Morocco
⁴Laboratory of Geophysics and Natural Risks, Department of Geology, Mohammed V University, Faculty of Sciences, Rabat, Morocco

Article history Received: 11-09-2020 Revised: 12-12-2020 Accepted: 16-12-2020

Corresponding Author: Dakir Ibrahim Laboratory of Environmental Geology and Natural Resources, Cartography and Digital Technologies Team, Department of Geology, Faculty of Science, Abdelmalek Essaadi University, Tetouan. Morocco Email: dakiribrahim@gmail_com **Abstract:** Electrical resistivity and Very Low Frequency electromagnetic profiles (VLF) were introduced for a litho-structural study in the locality of Douira, located in the Errachidia basin. The electrical reading was carried out with a Schlemberger type device. The VLF profiles were implemented with a different length reaching 1600 m and profile lines oriented in the direction NNW-SSE and N-S. The VLF measurements were obtained using an Iris T-VLF instrument and the data filtering was carried out using the KHFFILT software. The Karous-Hjelt filter was applied to the real component of the secondary electromagnetic field. The qualitative interpretation of the resistivity results has shown that the formations identified in the locality of Douira are more resistant than those encountered around Ain Al Atti, this difference is the salinity coming from marly formations with intercalation of gypsum deep in the vicinity by Ain Al Atti. The semi-quantitative interpretation of the VLF results showed the presence of a main fracture zone, oriented NW-SE, which feeds the Ain Al Atti source.

Keywords: Errachidia Basin, Electrical Survey, Fracture Zone, VLF-EM, Resistivity, Filtering KH, Groundwater Exploration

Introduction

In Morocco, the continual decrease in water resources is the current major problem, since water is a determining factor for sustainable development and improving the quality of life. The quantitative and qualitative management of groundwater and surface water resources is currently a necessity, given the continuous increase in water demand, on the one hand and the degradation of water quality observed during the past few years on the other hand. Water resources are limited and drought affects most of the country and becomes seriously threatening.

In the south-east of Morocco, or more precisely in the province of Errachidia, the scarcity and randomness of precipitation, the increasingly worrying desertification, the overexploitation of groundwater, linked to the strong demographic pressure and agriculture and animal husbandry activities have a great influence on the piezometric levels of groundwater (Amharref, 1991). The salt water from the source of Ain Al Atti has a negative impact on the environment by inducing direct chemical contamination on the aquifers in place. This therefore requires urgent intervention to curb the polluting flow from this source (Ammary, 2007). The problem addressed by this study concerns the potential groundwater in the locality of Douira north of the Ain Al Atti source. This article helps assess this potential by identifying fracturing zones. Faults represent a special target for hydrogeological exploration and can play an important role in the supply of groundwater (Reinhard, 2006). To achieve the objective of this article to highlight the litho-structural parameters, the working procedure will be based on electrical soundings, in order to determine the vertical variation of the resistivity obtained at the level of Douira and at the level of Ain Al



Atti. On the other hand, it will be based on the Very Low Frequency Electromagnetic technique (VLF-EM) in order to locate the alignment of the fractured areas.

Geography and Geological Setting

The study area is located about 50 km south of Errachidia and 20 km north of Erfoud Fig. 1. It is a part of cretaceous basin of Errachidia (South-eastern Morocco). This basin is characterized by a stratigraphic series that ranges from paleozoic to quaternary. In the outcrop Fig. 2, the basin is generally composed of carbonate deposits of Turonian (Choubert and Faure-Muret, 1960), sandstone and

sand with gypsum intercalation attributed to Infracenomanian and sand clay with gypsum and anhydrite of Senonian (Choubert, 1948). Locally, from drilling data, the Infracenomanian deposits overlie the Paleozoic (angular discordance). The quaternary is presented by alluvium and conglomerates. It shows varying thickness layers between 5 and 40 m (Amharref, 1991).

Lithostratigraphy

The stratigraphic log in Fig. 3 shows the formations constituting the basin range from the Paleozoic to the Quaternary.



Fig. 1: Geographic situation of the study area



Fig. 2: Geological section of the Errachidia basin by (Margat, 1977)

Dakir Ibrahim *et al.* / Current Research in Geoscience 2020, Volume 10: 16.26 DOI: 10.3844/ajgsp.2020.16.26

Age		Lith	Description	Thickness (m)	
Quaternairy			0.00	Lake limestones, silts and conglomerates	5-40
Secondary	Cretacous	Senonien		Red sandstone sands with intercaltion of marl and gypsum Clay sands Sandstone level with calcareous intercalation	150-600
				Lumachellic massive limestones Bed limestone with marl intercalation	60-150
		Cenomanien		Green manes anhydrite	40-50
		Infracenomanien		Red sandstone Sandy clays Anhydrites Fine sands	80-400
	Jurassic	Jurassic continent		Coarse sandstone Red clays Conglomerate	150-280
		Bathonien Bajocien		Dolimitic limstones Alternation of limstones and marls	100-320
		Aalenien		Alternation of marls and dolomitic limestones	100-320
		Toarcien		Fossiliferous marls	50-250
		Domerien	(More or less marly dolomitic limestone	150-200
	Trias	Trias	369769	Red marls Basal doleritic Anhydritic red marls	500
Paléozoïc Schist, quartzite and limestone					

Fig. 3: Synthetic stratigraphic log of the Cretaceous basin of Errachidia (Amharref, 1991)

Materials and Methods

Electrical Survey

The electrical sounding consists of the analysis of the apparent resistivity of the subsoil, measured by a symmetrical quadrupole device (AMNB) for a succession of spacing of the electrodes AB. The investigation depth is adjusted by varying the distance between the current injection electrodes A and B. The difference in the potential measured between M and N reflects an apparent resistivity linked to the true resistivity and to the thicknesses of all the layers interested in the AMNB device. This method is widely used and generally provides a good initial approach Fig. 4. The apparent resistivity ρ_a is calculated from the current I and the potential difference ΔV (Equation 1). The coefficient K is called geometric factor. For Schlemberger configuration, the factor K can be calculated from the electrode spacing by (Equation 2):

$$\rho_2 = K \cdot \frac{\Delta V_{MN}}{I_{AB}} \tag{1}$$

$$K = 2\pi a \tag{2}$$

VLF Electromagnetic Survey

The Very Low Frequency Electromagnetic (VLF-EM) is based on the use of radio waves in the range of 15 to 30 kHz. The signal primary magnetic field (H_p) emitted by the VLF stations, can be captured in the field by the VLF instruments. When a conductor (e.g., a fracture zone) is crossed by the H_p electromagnetic field, an induced current (Current of Foucault) flows through it and produces a secondary magnetic field (H_s) out-of-phase with primary magnetic field oriented in any

direction (McNeill and Labson, 1991). In this case the conductive body acts as a second source (Kaya *et al.*, 2007). The resulting field from the sum of primary magnetic field and secondary magnetic field is elliptically polarized. This ellipse of polarisation has two components with the same frequency, but different amplitude and phase, (Eze *et al.*, 2004). The in-phase H_p is the Real component (R_e) proportional to the tilt τ (inclination of the major axis of the ellipse), while the out of phase H_p is the Imaginary component (Im) proportional to the ellipticity ε (the ratio between the minor and the major axis b/a). These two components τ and ε are described by the Equations (3 and 4) below (Saydam, 1981):

$$\tau = R_e / H_p \tag{3}$$

$$\varepsilon = I_m / H_p \tag{4}$$

During our study, the survey was carried out using the Receiver T-VLF Iris Instruments, operating in tilt angle mode, in order to measure the parameters of the ellipse of polarization, which are the tilt τ and the ellipticity ε Fig. 5. In this mode, it is convenient to operate with a transmitter (VLF station) which is located in the supposed strike ($\pm 45^{\circ}$) of the prospected target for a maximum coupling. For detecting the supposed fractures in the study area, the GBR station located in Rugby (England) has been chosen, with a power of 750 KW, which emits a signal with a frequency of 16 KHz. On the fieldwork, 6 VLF EM profiles were conducted, with profile length reaches 1600 m. Readings were taken respecting a spacing of 20 m. The profiles lines were oriented NNW-SSE and N-S directions.



Fig. 4: Representation of the schlumberger type device



Fig. 5: Field components of VLF field from transmitter at Remote Distance (Ezepue, 1984)

For our measurements, in one hand, the Karous and Hjelt (KH) filter has been applied to the real component. This filter permits the draw of apparent current density cross-sections, which show the response of the conductor in depth (Karous and Hielt, 1983). Qualitatively, it is possible to discriminate between conductive and resistive structures using apparent current density cross-section (Karous and Hielt, 1977), where a high positive value corresponds to conductive structure and low negative values are related to resistive one (Benson et al., 1997; Sharma and Baranwal, 2005), In the other hand, the Fraser filter has been applied to the real component too and presented in the form of a contour map. Therefore, the filtered real component will always show a positive peak above an anomalous zone (Fraser, 1969). In order to perform Karous-Hjelt and Fraser filtering on VLF EM data, the software KHFFILT is used in the interpretation of the measurements along VLF traverses.

Results

In order to obtain better recognition of the study site, we used two geophysical techniques; the method of electrical and electromagnetic exploration. In fact, two Schlumberger-type electrical soundings were carried out (AB = 600 m long) and six electromagnetic profiles (Fig. 6).

Electrical Prospecting

The electrical survey N°1 is carried out next to the village of Douira on the left bank of the national road

 $N^{\circ}13$ (about 100 m) which leads to Erfoud. Its direction is N-S (Fig. 7a). While the survey $N^{\circ}2$ was carried out at 250 m from the source of Ain El Atti in direction NW-SE Fig. 7b.

Electromagnetic Prospecting

We focus on the analysis of positive Karous-Hjelt anomalies, for semi-quantitative interpretation and visualization of the target. For the data filtered by KH, the cross-sectional curves of the apparent current density were produced (Figs. 9 to 11). Here, the traces KH are represented for three types of lines: 2, 3 and 4.

Discussion

The survey in Fig. 6 suggests the presence of a conglomerate formation with gravel on the surface, of resistivity $\rho = 2000$ Ohm.m and of thickness 1 m. The latter rests on a formation of weathered sandstone with clays, of resistivity $\rho = 300$ Ohm-m and thickness 7 m, which overcomes a formation of limestone's of resistivity $\rho = 3060$ Ohm-m and thickness 64 m. In depth, we find sandy sandstones with clays of resistivity $\rho = 750$ Ohm-m (Fig. 8a). For the survey in Fig. 7b highlights the sands and gravels in surface (resistivity $\rho = 300$ Ohm-m and thickness 2.6 m) overcome a claymarly formation with resistivity $\rho = 9$ Ohm-m and thickness 16 m, which rests on the sandstones ($\rho = 270$ Ohm-m and thickness 16 m). At the base, we find a marl formation of resistivity 8 Ohm-m (Fig. 8b).

Dakir Ibrahim et al. / Current Research in Geoscience 2020, Volume 10: 16.26 DOI: 10.3844/ajgsp.2020.16.26



Fig. 6: Location of electrical and electromagnetic surveys



Fig. 7: Electrical survey: a- Douira, b- Ain Al Atti

Dakir Ibrahim et al. / Current Research in Geoscience 2020, Volume 10: 16.26 DOI: 10.3844/ajgsp.2020.16.26



Fig. 8: Synthetic electric logs: a- Survey carried out in Douira, b- Survey carried out in Ain Al Atti

Taking into account the electrical logs (Fig. 8a and 8b), we observe that for the upper part of the field investigated, the formations encountered at Douira are relatively the same encountered at Ain Al Atti, however in depth we find a variation of facies (gray and limestone at the level of Douira and the marls at the level of Ain Al Atti), these marly formations with intercalation of gypsum encountered in depth, caused a significant decrease in the electrical resistivity in the vicinity of Ain Al Atti.

The line (Fig. 8a) represents the variation of the apparent section of the current density as a function of the distance, along line 2. The plot shows a prominent positive response between 900-1000 m, from the start of the profile, corresponding to a conductive ax, resulting in a fracture zone at a depth ranging from the surface to 170 m. The results of line 3 are represented as a cross section of apparent current density in Fig. 9a.

The plot reveals a positive anomaly linked to the presence of a conductive target at 1580 m from the start of the profile. It is interpreted as a fracture zone at a depth ranging from surface to 140 m. the KH Filtering data from the rest of these lines does not show any pronounced anomalies. The result of line 4 (Fig. 10a) suggests the presence of a conductive anomaly at a distance of 350 m which corresponds to the same fractured zone. The filtered Fraser data reveals the presence of anomalies which are located and correlated to the anomalies obtained by Karous Hjelt filtration. Three positive anomalies were distinguished, where the apex of the anomalies was observed at locations L2 (Fig. 8b), L3 (Fig. 9b) and L4 (Fig. 10b). These are linked to the presence of conductive structures, interpreted as fracture zones. This fracture has a major NW-SE direction which confirms the work carried out at the level of the locality Zaouia Jdida by Ait (Bahammou et al., 2019) is schematized on the model of Fig. 12.







Fig. 10: Electromagnetic VLF-L3: A- apparent current density cross section, b- fraser filtering



Fig. 11: Electromagnetic VLF-L4: a- apparent curent density cross section, b- fraser filtering



Fig. 12: Schematic model of the location of the fracturing zone

Conclusion

In this article, we have contributed to the litho-structural study at the level of the locality of Douira and near the artesian source of Ain Al Atti, by combining the electrical technique and the very low frequency electromagnetic method. The results of the electrical soundings showed the presence of very conductive grounds in the vicinity of the source which is due to the capillary rise of the salinity coming from marly formations with gypsum intercalations in depth. Data filtered by VLF revealed the presence of a fracture anomaly which reached 150 m in depth and which is oriented NW-SE. The identified fracture anomaly constitutes the potential groundwater circulation area that led to the supply of the Ain Al Atti source.

Given the importance of the results obtained at the level of the prospected area, it is advisable to carry out other electromagnetic profiles and electrical trails in the northern part of the locality of Douira and to carry out a mechanical survey at the level of these conductive anomalies. To block the loss of salt water from the Ain Al Atti source.

Acknowledgement

The authors are sincerly grateful to the Faculty of Science and Technology Errachidia (FSTE) for proveding the geophysical machines used in this study.

Author's Contributions

Dakir Ibrahim: I carried out the geophysical measurements, I contributed to the interpretation of the data, the writing of the article.

Benamara Ahmed: I am the professional supervisor of Mr. DAKIR. I helped the main author with the interpretation of geophysical data.

Aassoumi Habiba: I am the administrative supervisor of Mr. DAKIR. Helped the main author with manuscript writing and translation.

Ouallali Abdessalam: I contributed to the field level; we carried out together several geophysical companies.

Ait Bahammou Youssef: I contributed to the field level; we carried out together several geophysical companies.

Ethics

This article is original and contains unpublished data. The corresponding author confirms that all of the other authors have read and approved the manuscript and no ethical issues involved

Data Availability

The data used to support the findings of this study are available from the corresponding author upon request.

References

- Amharref, M. (1991). Contribution à l'étude hydrogéologique de la vallée du Ziz (province d'Errachidia Sud-Est du Maroc): Incidences respectives de la sécheresse et du Barrage Hassan Addakhil sur les ressources à l'aval (Doctoral dissertation, Besançon).
- Ammary, B. (2007). Etude géochimique et isotopique des principaux aquifères du bassin Crétacé d'Errachidia et de la plaine de Tafilalet.
- Bahammou, Y. A., Benamara, A., Ammar, A., & Dakir, I. (2019). Fracture zones detection for groundwater exploration integrating Resistivity Profiling and Very Low Frequency electromagnetic methods (Errachidia basin, Morocco). Contributions to Geophysics and Geodesy, 49(2), 181-194.
- Benson, A. K., Payne, K. L., & Stubben, M. A. (1997). Mapping groundwater contamination using dc resistivity and VLF geophysical methods–A case study. Geophysics, 62(1), 80-86.
- Choubert, G. (1948). Essai sur la paléogéographie du Mésocrétacé marocain. Volume jubilaire de la Société des Sciences Naturelles du Maroc, 1920210(1945), 307-329.
- Choubert, G., & Faure-Muret, A. (1960). Evolution du domaine atlasique marocain depuis les temps paléozoïques. Mémoire hors-série-Société géologique de France, (1), 447-527.
- Eze, C. L., Mamah, L. I., & Israel-Cookey, C. (2004). Very low frequency electromagnetic (VLF-EM) response from a lead sulphide lode in the Abakaliki lead/zinc field, Nigeria. International journal of applied earth observation and geoinformation, 5(2), 159-163.
- Ezepue, M. C. (1984). The geologic setting of lead-zinc deposits at Ishiagu, southeastern Nigeria. Journal of African Earth Sciences (1983), 2(2), 97-101.
- Fraser, D. C. (1969). Contouring of VLF-EM data. Geophysics, 34(6), 958-967.
- Karous, M., & Hjelt, S. E. (1977). Determination of apparent current density from VLF measurements: Report. Department of Geophysics, University of Oulu, Finland, Contribution, 89, 19.
- Karous, M., & Hjelt, S. E. (1983). Linear filtering of VLF dip-angle measurements. Geophysical prospecting, 31(5), 782-794.
- Kaya, M. A., Özürlan, G., & Şengül, E. (2007). Delineation of soil and groundwater contamination using geophysical methods at a waste disposal site in Çanakkale, Turkey. Environmental monitoring and assessment, 135(1-3), 441-446.
- Margat, J. (1977). Etude hydrogéologique du bassin Quaternaire de Tafilalet. Ressources en Eau du Maroc, p310-380.

- McNeill, J. D., & Labson, V. F. (1991). Geological mapping using VLF radio fields. In Electromagnetic Methods in Applied Geophysics: Volume 2, Application, Parts A and B (pp. 521-640). Society of Exploration Geophysicists.
- Reinhard, K. (Ed.). (2006). Groundwater Geophysics: A Tool for Hidrogeology. Springer.
- Saydam, A. S. (1981). Very low-frequency electromagnetic interpretation using tilt angle and ellipticity measurements. Geophysics, 46(11), 1594-1605.
- Sharma, S. P., & Baranwal, V. C. (2005). Delineation of groundwater-bearing fracture zones in a hard rock area integrating very low frequency electromagnetic and resistivity data. Journal of Applied geophysics, 57(2), 155-166.